

## Lecture Ch. 12a

- Review of simplified climate model
- Current research
  - Aerosols, precipitation, and evaporation
- Course evaluations

Curry and Webster, Ch. 12  
 For Tuesday: Read Ch. 13  
 For Dec 6: Review Past Homework, Quizzes, Reading, Midterm

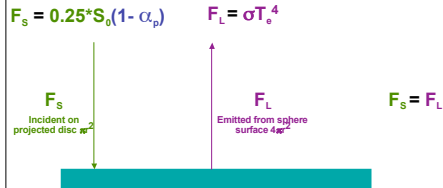
## Simplified Climate Model

- Atmosphere described as one layer
  - Albedo  $\alpha_p \sim 0.31$ : reflectance by surface, clouds, aerosols, gases
  - Shortwave flux absorbed at surface  $F_s = 0.25 \cdot S_0 (1 - \alpha_p)$
- Earth behaves as a black body
  - Temperature  $T_e$ : equivalent black-body temperature of earth
  - Longwave flux emitted from surface  $F_L = \sigma T_e^4$

Curry and Webster, Ch. 12 pp. 331-337; also Liou, 1992

## Simplified Climate Model

- Incoming shortwave = Outgoing longwave
- Energy absorbed = Energy emitted

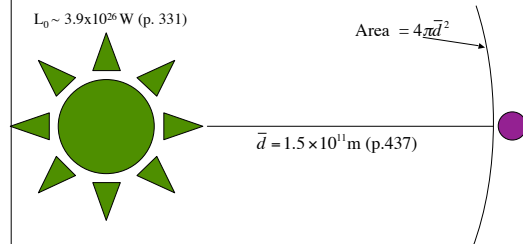


## Solar Constant

- Luminosity of the sun
- Irradiance at earth  $S_0 = L_0 / (4\pi d^2) = 1.4 \times 10^3 \text{ W/m}^2$

$L_0 \sim 3.9 \times 10^{26} \text{ W}$  (p. 331)

Area =  $4\pi d^2$



## Simplified Climate Model

- At thermal equilibrium (why?)

$$F_s = F_L$$

$$0.25 \cdot S_0 (1 - \alpha_p) = \sigma T_e^4$$

$$T_e = [0.25 \cdot S_0 (1 - \alpha_p) / \sigma]^{0.25}$$

$$T_e \sim 255 \text{ K}$$

- Observed surface temperature  $T = 288 \text{ K}$
- What's missing?

## Sensitivity to Albedo

- What if albedo changes?

$$T_e = [0.25 \cdot S_0 (1 - \alpha_p) / \sigma]^{0.25}$$

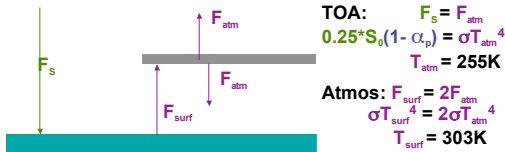
$$\alpha_p = 0.31, T_e \sim 255 \text{ K}$$

$$\alpha_p = 0.30, T_e \sim ?$$

- 1% decrease in albedo warms temperature 1K
- 1% increase in albedo cools temperature 1K

## Add an Atmosphere!

- Atmosphere is transparent to non-reflected portion of the solar beam
- Atmosphere in radiative equilibrium with surface
- Atmosphere absorbs all the IR emission



## What's wrong?

- With no atmosphere,  $T_{surf} = 255K$
- With "atmosphere",  $T_{surf} = 303K$
- From observations,  $T_{surf} = 288K$
- Real atmosphere:
  - Not perfectly transparent to incoming solar
  - Not perfectly opaque to infrared
  - Not in pure radiative equilibrium with surface
- Three assumptions were wrong!

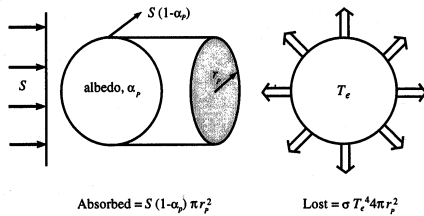
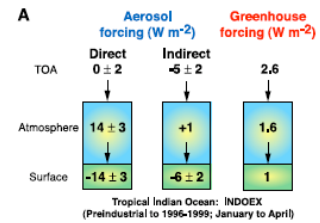
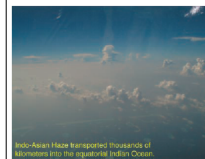


Figure 12.1 Radiative energy balance of the Earth. The Earth absorbs energy from the sun and loses energy through longwave emission.

## Current Research: Radiation

- Large surface cooling:  $-20 W m^{-2}$ 
  - Reduces evaporation, short-circuiting precipitation

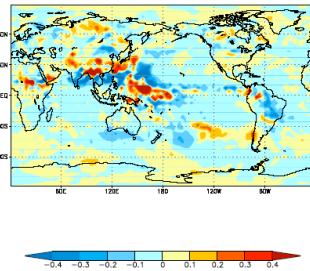


Ramanathan et al., 2001

## Current Research: Global Models

- Aerosol impacts on rain are not local

Change in JJA mean precipitation (mm day<sup>-1</sup>) between the 6-year perturbation and the 12-year control.



Erlick, Ramaswamy, and Russell, 2005

Vol 439/30 June 2005/doi:10.1038/nature03671

nature

## PROGRESS

Strong present-day aerosol cooling implies a hot future

Meinrat O. Andreae<sup>1</sup>, Chris D. Jones<sup>2</sup> & Peter M. Cox<sup>3</sup>

"...aerosols counteract warming by an uncertain, but potentially large, amount... Strong aerosol cooling would imply that future global warming may proceed at or even above the upper extreme of the range projected by the IPCC."

