

Oceanic phytoplankton, atmospheric sulfur, cloud albedo and climate

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Presented By

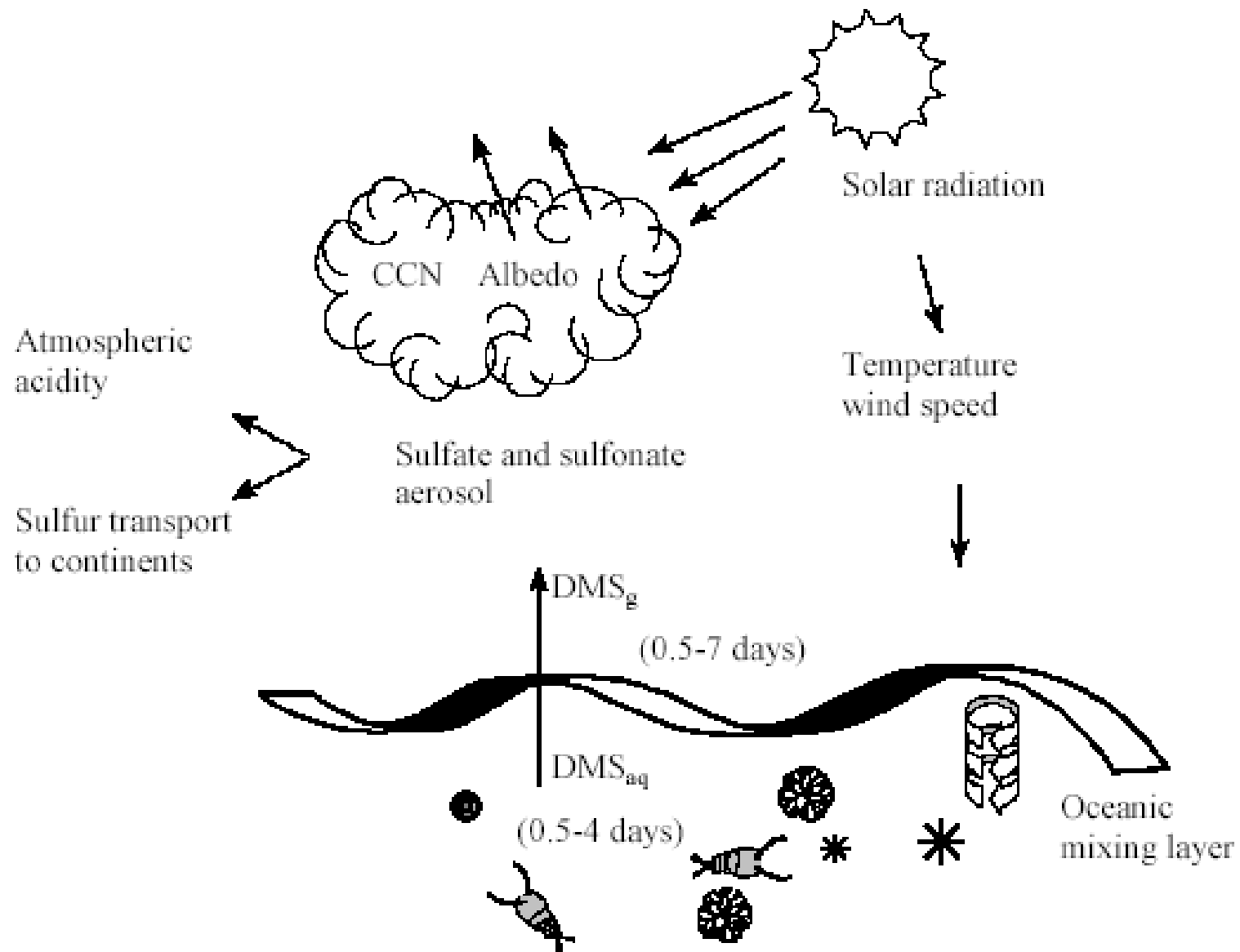
John Holecek and Aihua Zhu

Major significance

- Proposal that biota could regulate climate
 - Hypothesized possible mechanisms
 - More data was needed to support theory
- Stimulated a great deal of scientific interest
- Cited 1203 times!
- Many assumptions made

What we knew before

- Radiation—Twomey 1977, indirect effect
- Gaia hypothesis – Lovelock 1974
- Aerosols (SO_4^{2-}) as CCN
- Global sulfur budget – Andreae 1985



Global Sulfur budget

TABLE 6. Sulfur emissions from natural and anthropogenic sources expressed in 10^9 mol/a

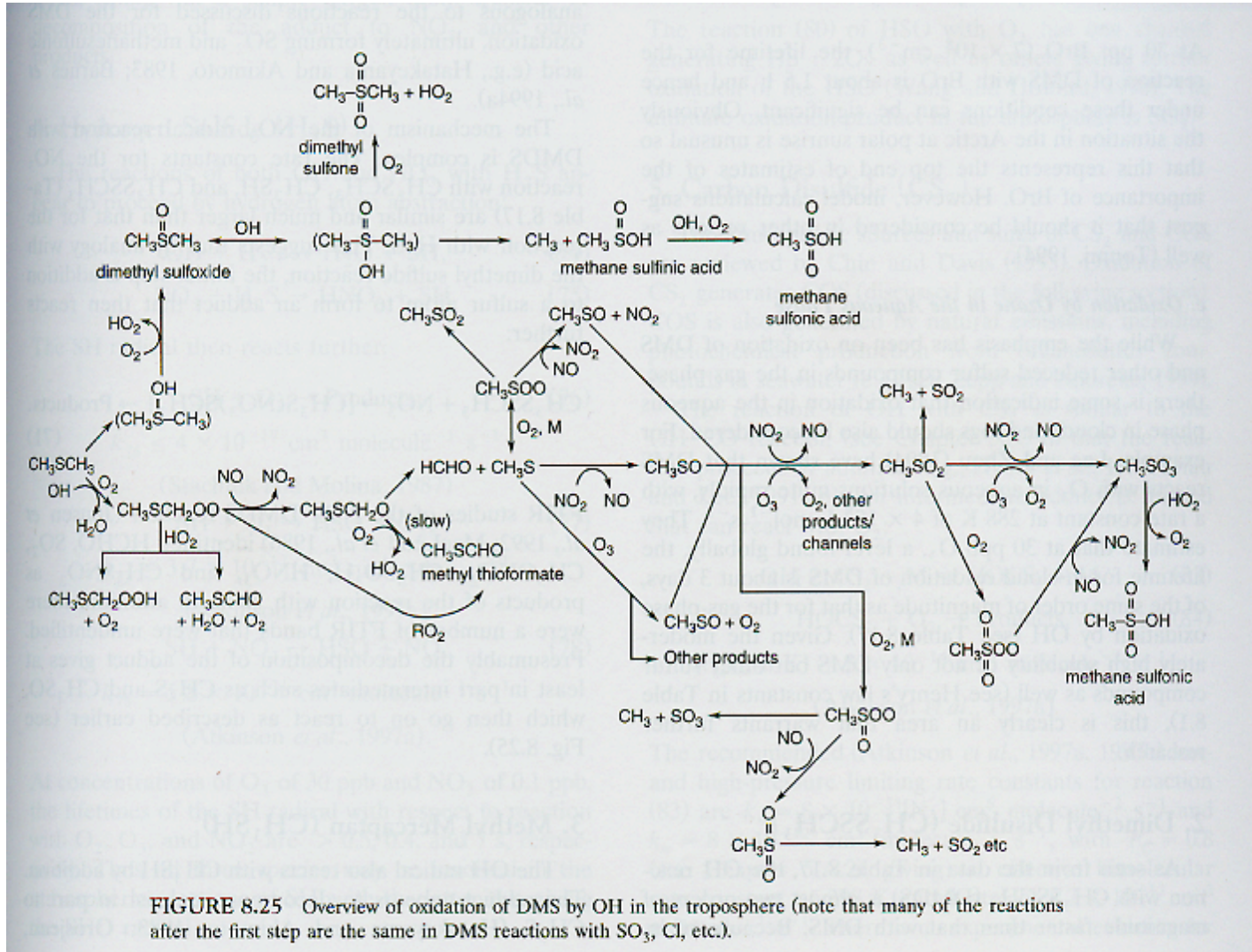
Region	Oceanic	Terrestrial	Volcanic	biomass burning	Anthropogenic	Biogenic/total ¹	Natural/total ²
80°-65°N	4	0.02	2	0.4	3	40%	62%
65°-50°N	19	0.40	43	2.3	534	3	10
50°-35°N	31	0.95	53	3.3	942	3	8
35°-20°N	46	2.05	37	7.1	598	7	12
20°-5°N	79	2.52	54	20.7	106	31	52
5°N-0°	26	1.14	17	4.2	18	41	67
0°-5°S	25	1.10	27	3.6	16	36	73
5°-20°S	82	2.11	45	17.3	47	44	67
20°-35°S	60	0.86	2	9.2	153	27	28
35°-50°S	60	0.08	8	0.7	24	65	73
50°-65°S	50	0.00	0	0.2	1	98	98
65°-80°S	4	0.00	1	0.0	0	79	100
N. Hemisphere	200	7.1	210	38	2200	8	16
S. Hemisphere	280	4.1	83	31	240	45	58
Global	480	11.	290	69	2400	15	24

¹ (Marine + Terrestrial)*100/(Marine + Terrestrial + Volcanic + Anthropogenic + Biomass burning).

² (Marine + Terrestrial + Volcanic)*100/(Marine + Terrestrial + Volcanic + Anthropogenic +

Bates, T. S., Lamb, B.K., Guenther, A., Dignon, J., Stoiber, R.E. (1992). "Sulfur emissions to the atmosphere from natural sources." J. Atmos. Chem. **14**: 315-327

DMS oxidation by OH

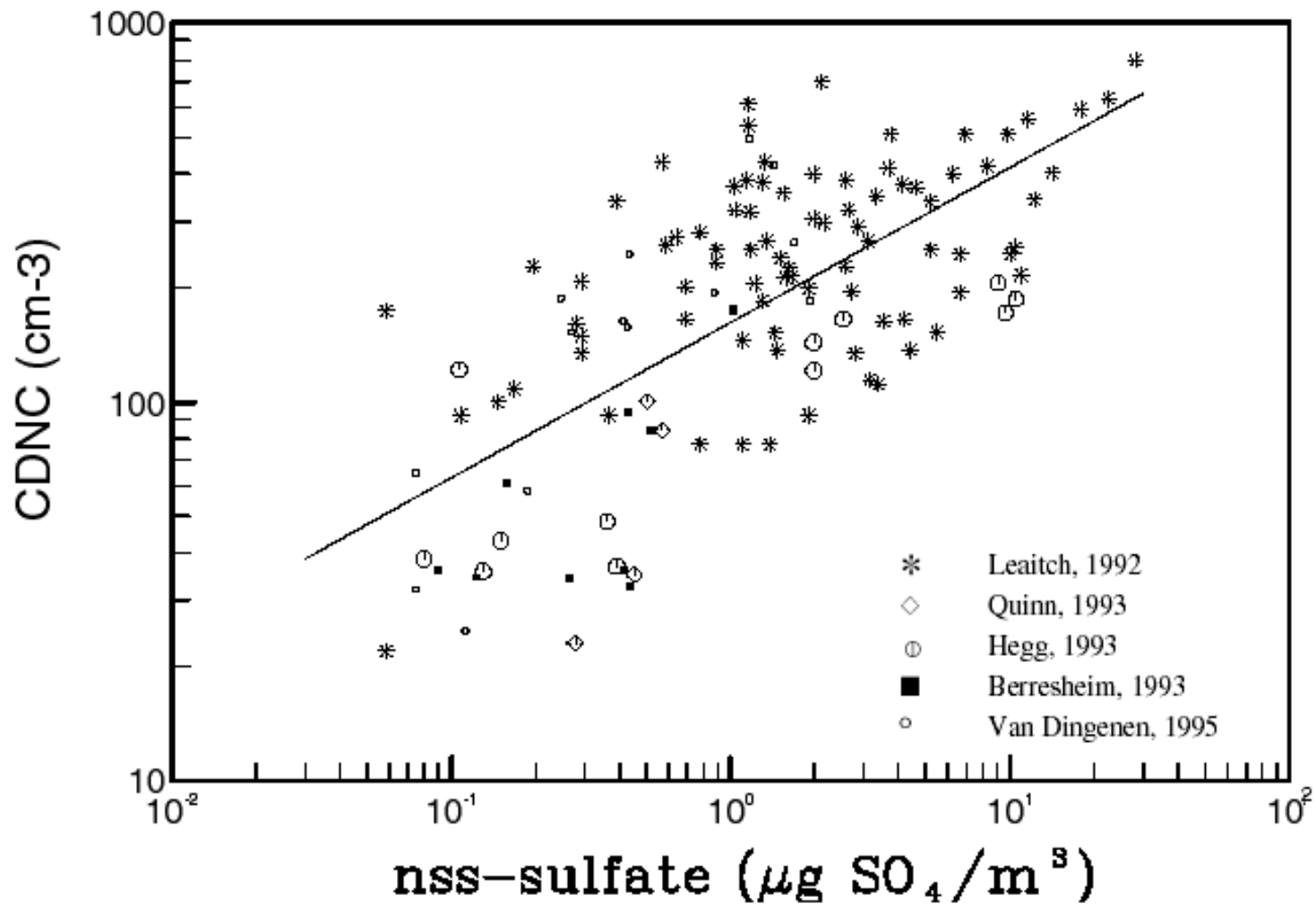


NSS-SO₄²⁻ particles are the main contributor to CCN

- 1) Significant fraction of submicrometer particles are active CCN
- 2) Most CCN are composed of water soluble materials
- 3) The total number-population of NSS-SO₄²⁻ agrees with measured CCN population
- 4) Much of the light-scattering aerosol in marine air is volatile at elevated temperatures
- 5) The turnover time of CCN is the same order as NSS-SO₄²⁻

CLOUD DROPLET NUMBER CONCENTRATION

Dependence on Non-Seasalt Sulfate



Boucher and Lohmann, 1995

Potential effects of NSS-SO₄²⁻ variations on cloud properties

Basic

$$L = (4/3) \pi r^3 \rho N$$

equation:

Three variables: L (liquid water content, gm⁻³)

N (number-density of droplets, m⁻³)

r (droplet radius, m)

1) Hold r fixed: T↑ → L↑ → N↑ → ↑

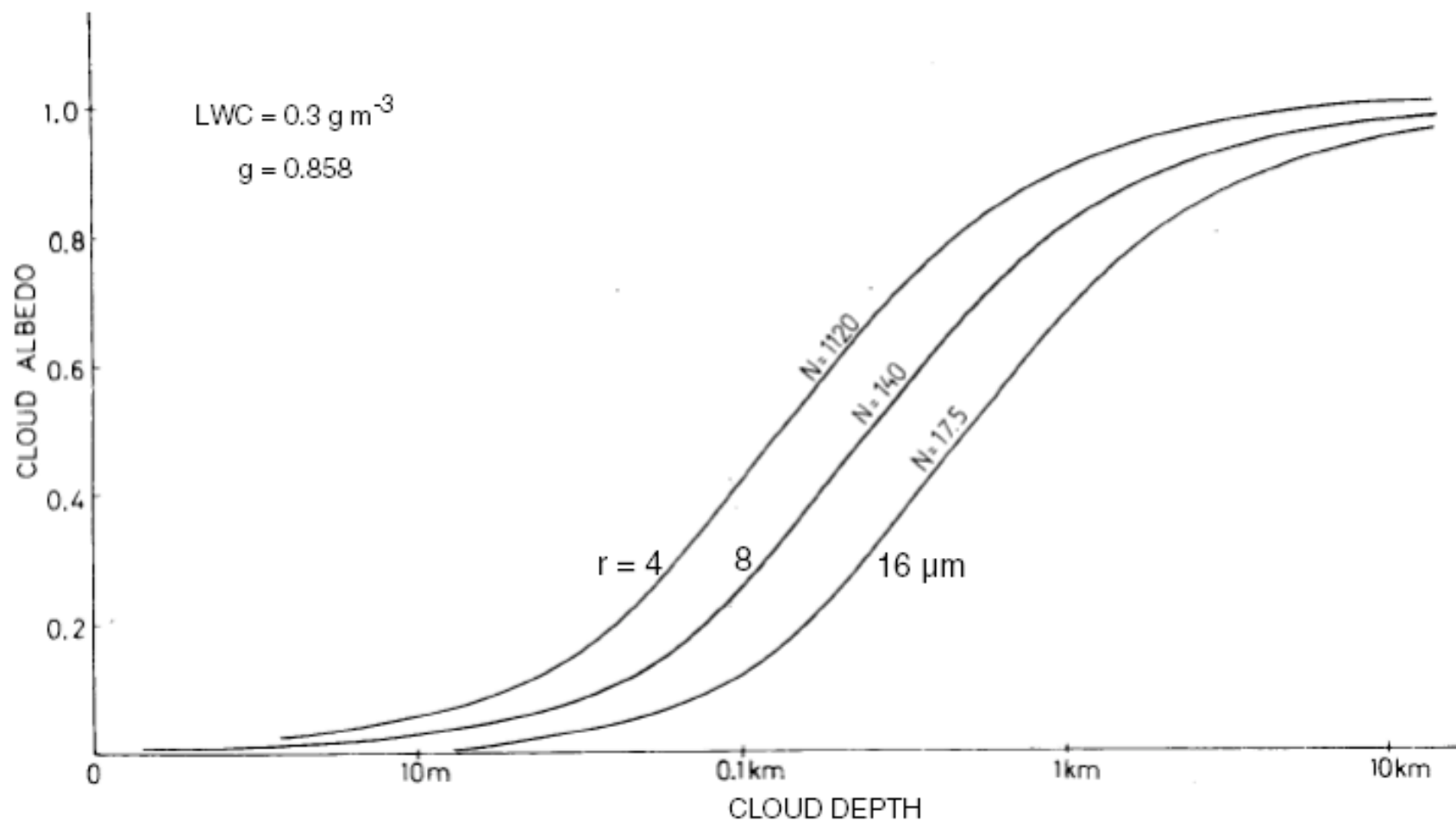
2) Hold N fixed: T↑ → L↑ → r↑ → ↓

3) Hold L fixed: T↑ → DMS↑ → N↑ → r↓ → ↑

**----Twomey's
effect**

DEPENDENCE OF CLOUD ALBEDO ON CLOUD DEPTH

Influence of Cloud Drop Radius and Concentration



Twomey, *Atmospheric Aerosols*, 1977

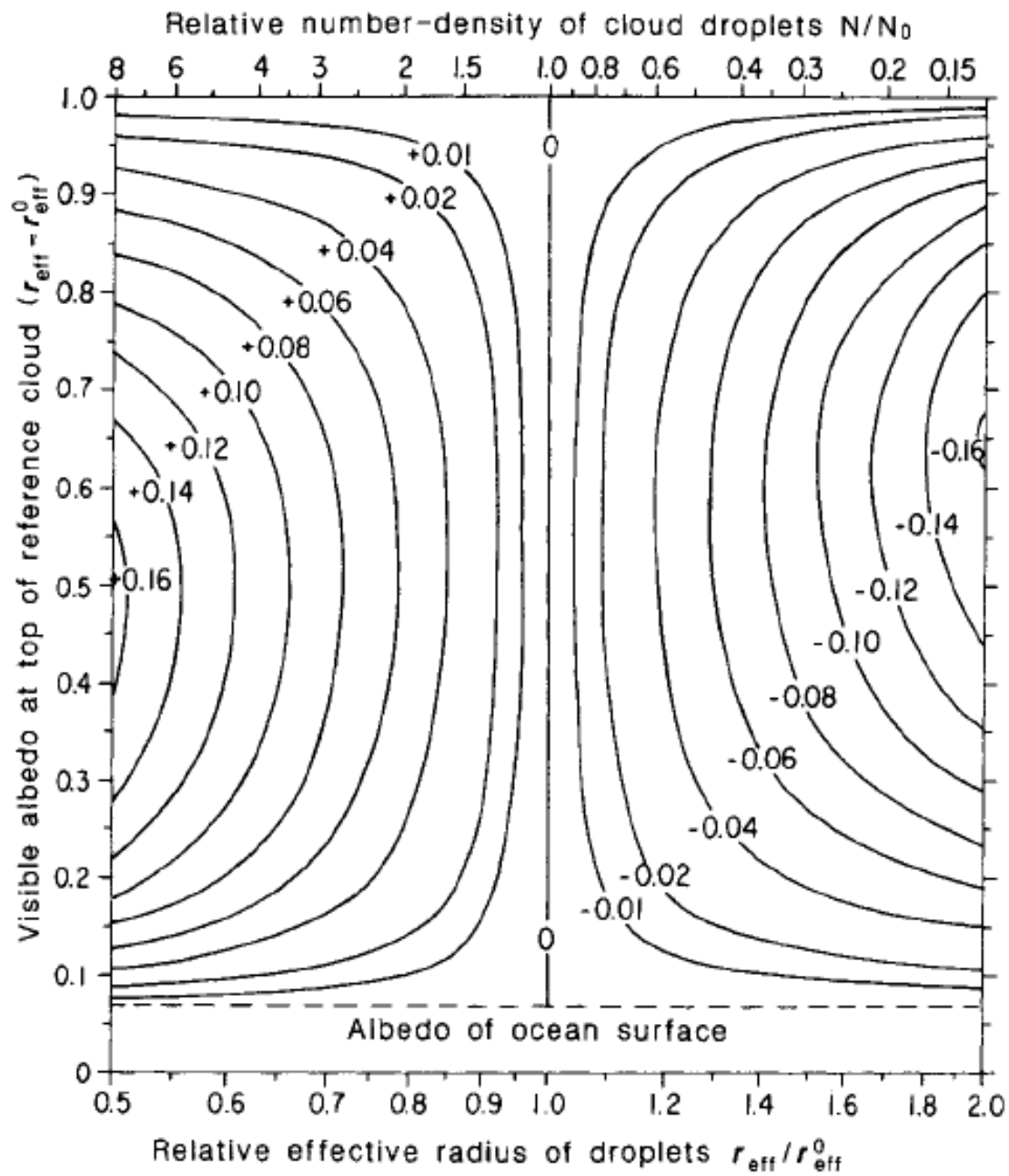


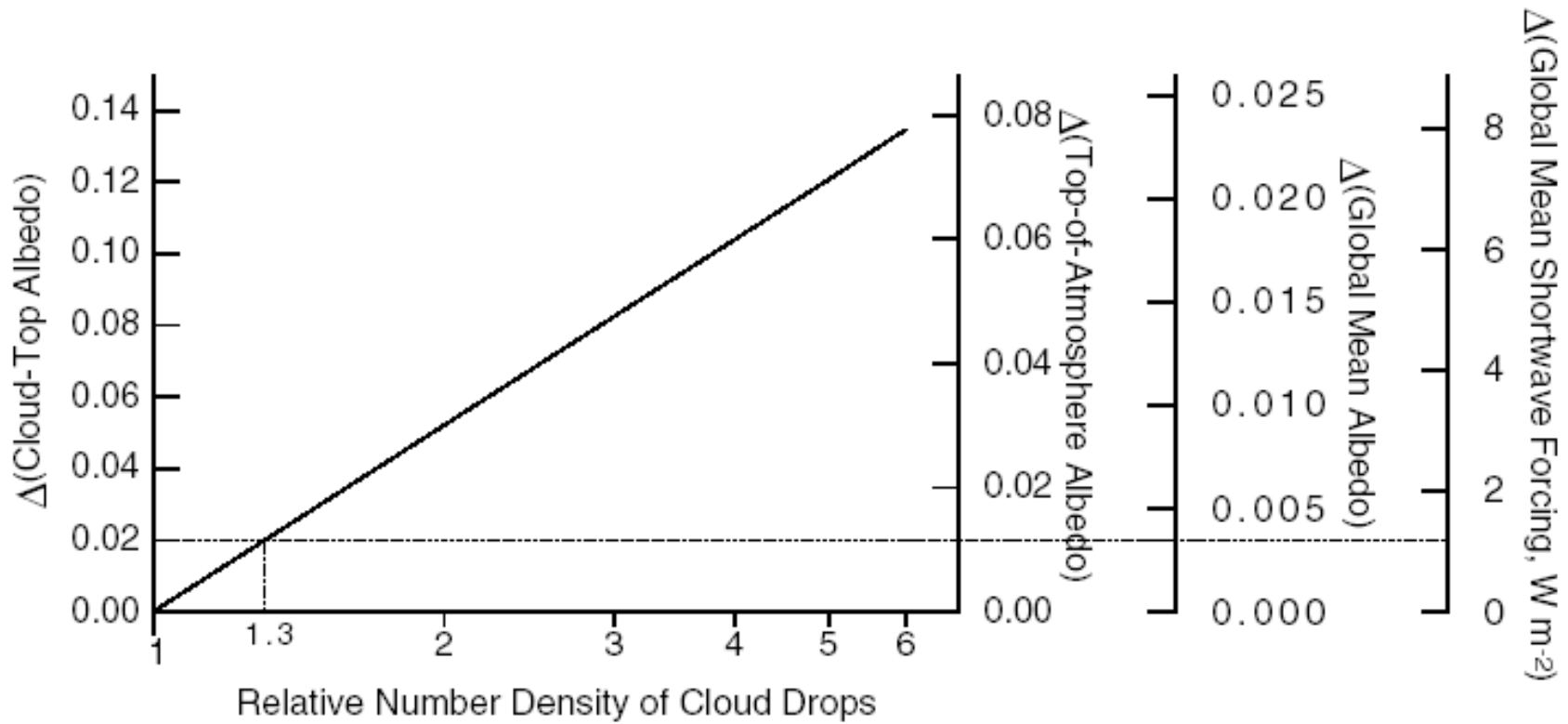
Figure
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Table 1

Table 1 Climatic effect caused by increasing CCN concentration over the ocean

<i>a</i> Global annual average cloud cover (ocean areas only)			<i>b</i> Example: effect on surface climate due to increasing CCN concentration N by 30% while holding liquid water path fixed		
Cloud type*	Ocean area covered (%)	Earth covered by oceanic clouds (%)		For area covered by oceanic stratiform water clouds	Averaged over Earth's surface area
Non-overlapped St/Sc†	25.2	17.6			
Non-overlapped As/Ac‡	10.8	7.5	Imposed change in N	+30%	
As/Ac overlapped with St/Sc§	8.8	6.1	Change in r_{eff}	-10%	
Nimbostratus, cumulus, cumulonimbus	not applicable (optically thick; high albedo)		Change in 0.5-0.7- μm albedo at TOC #	+0.02	
Cirrus	not applicable (ice)		Change in 0.5-0.7- μm albedo at TOA**	+0.018	
Total cover of oceanic stratiform water clouds (As/Ac+St/Sc) not overlapped with cumuliform clouds	44.8¶	31.2	Change in solar albedo at TOA**	+0.016	+0.005
			Equivalent change in solar constant††		-0.7%
			Change in global-average surface temperature‡‡		-1.3 K

SENSITIVITY OF ALBEDO AND FORCING TO CLOUD DROP CONCENTRATION



Schwartz and Slingo (1996)

Global climate and DMS emission

A stable negative

feedback

Sea-to-air mass flux of

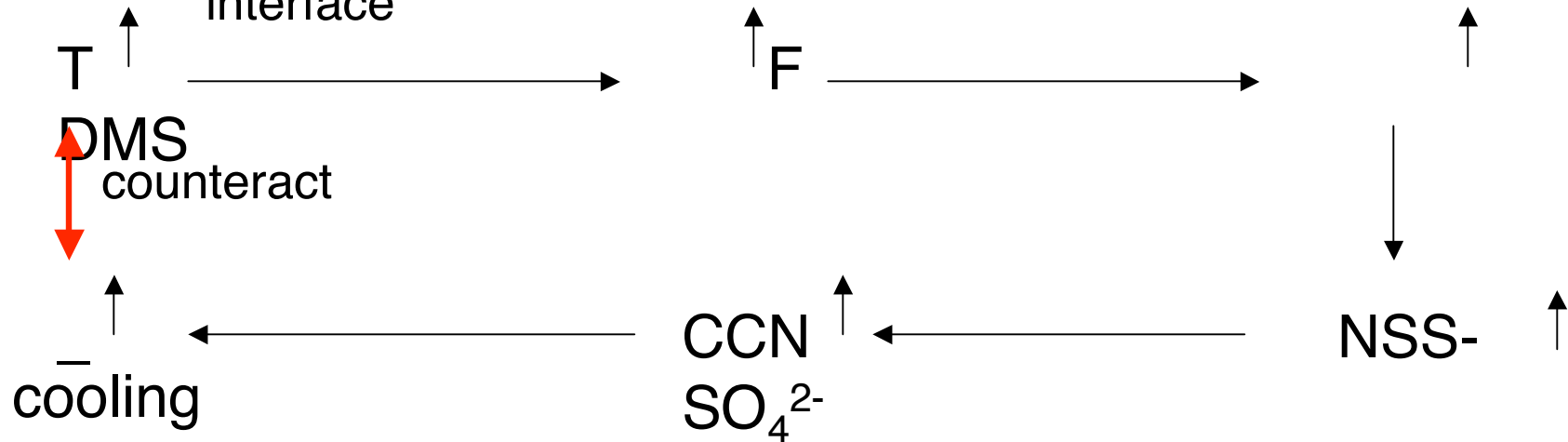
$$F = A \cdot k \cdot \Delta c$$

DMS:

A: total ocean surface $\downarrow T \rightarrow \downarrow A \rightarrow \downarrow F$

K: transfer velocity

Δc : concentration gradient across the air/sea interface



Climatic Feedback Loop

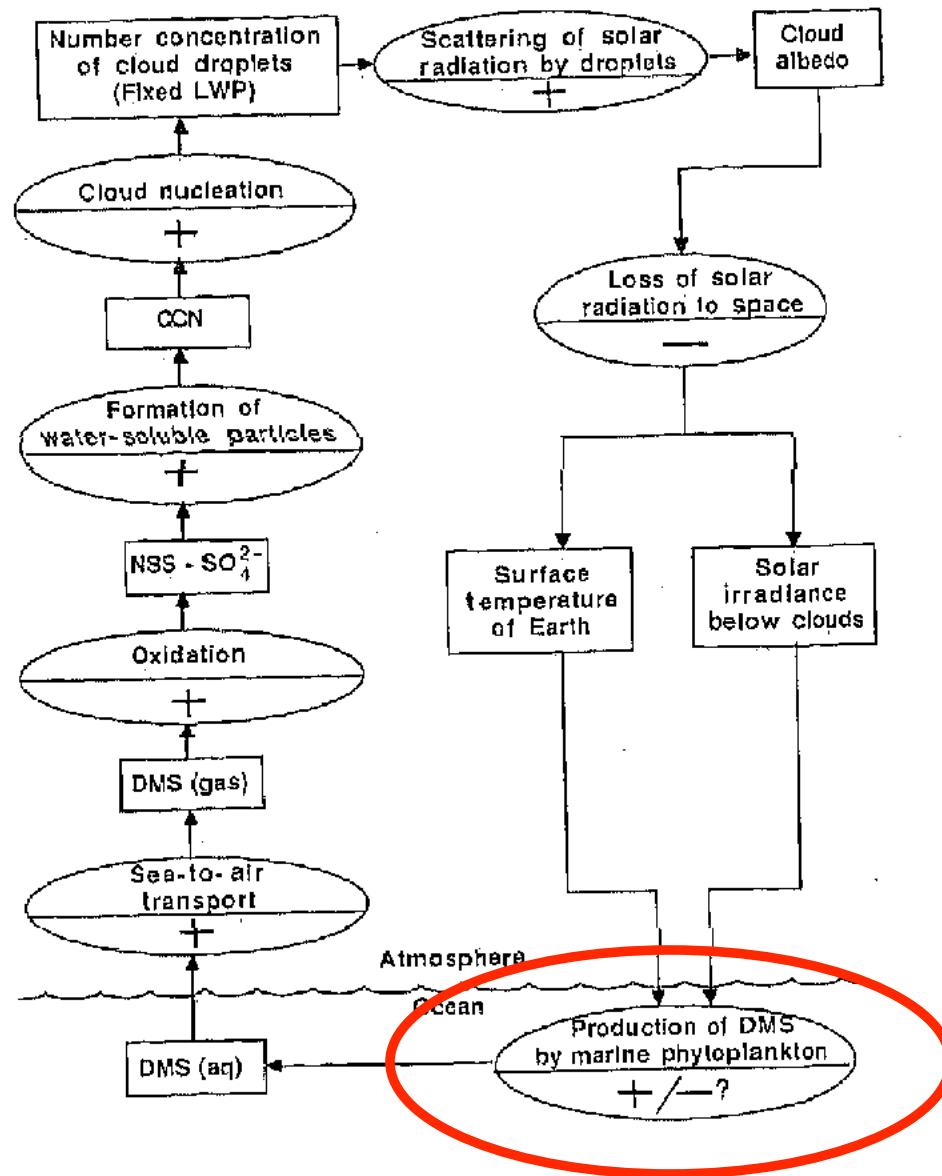


Figure 2

Atmospheric DMS concentrations

- Independent of the rate of primary production, the warmest, most saline, and most intensely illuminated regions of the oceans have the highest rate of DMS emission to the atmosphere
- Key fact when considering possible climatic feedback mechanisms

Gaia theory*

- Early after life began it acquired control of the planetary environment
- This homeostasis by and for the biosphere has persisted until the present
- Why would phytoplankton evolve to produce DMS?
 - Many theories for the mechanism, little supporting data
 - Reaction to salt stress

*Lovelock, J. E. and L. Margulis (1974). "Atmospheric homeostasis by and for the biosphere. Gaia hypothesis." Tellus **26**(1-2): 2-10.

Future work

- Intense study in many areas in attempts to support or disprove the theory
 - Biological role of DMSP & possible explanations for its evolutionary development Simo, R. (2001). "Production of atmospheric sulfur by oceanic plankton: biogeochemical, ecological and evolutionary links." Trends in Ecology & Evolution **16**(6): 287-294.
 - Understanding ocean-atmosphere DMS fluxes Kettle, A. J. and M. O. Andreae (2000). "Flux of dimethylsulfide from the oceans: A comparison of updated data sets and flux models." Journal of Geophysical Research, [Atmospheres] **105**(D22): 26793-26808.
 - Connection between DMS and CCN Pandis, S. N., L. M. Russell, et al. (1994). "The relationship between DMS flux and CCN concentration in remote marine regions." Journal of Geophysical Research, [Atmospheres] **99**(D8): 16945-57. [Reasonable agreement between observations and model](#)